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The Tectonic Evolution of the Central Andean Plateau and Geodynamic Implications for the Growth of Plateaus

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Geologic Setting



- Altiplano: internally drained with low relief topography and an average modern elevation of ~4km
- Altiplano region:
 Overlies a zone of
 normal subduction
 (~30° dip) of Nazca
 plate between two
 zones of flat slab
 subduction

Key Questions in the Central Andes

- Geodynamic processes for building the Andean plateau?
 - Long-term crustal shortening/thickening causes gradual isostatic surface uplift
 - Isostatic surface uplift by magmatic addition
 - Surface uplift by delamination/convective removal of dense lower lithosphere
 - Redistribution of crustal material by mid to lower crustal flow



Altiplano basin evolution



Western Cordillera



Eastern Cordillera – Salla Formation

End Member Model I

Slow and Gradual Surface Uplift

Surface uplift rates = 0 to 0.25 mm/yr

Crustal thickening/continuous subduction of foreland lithosphere



End Member Models II

Rapid Pulses of Surface Uplift

Surface uplift rates >0.4 mm/yr

Removal of lower lithosphere

Lower crustal flow



End Member Model II



Rapid Surface Uplift

- Decoupled deformation and surface uplift
- Proposed Mechanisms: Removal of dense lower lithosphere and/or lower crustal flow



Large magnitude surface uplift of the central Altiplano between 10 and 6 Ma.

Garzione et al. (2006, EPSL; 2008, Science); Ghosh et al. (2006, Science)

⁾ Paleobotany: Gregory-Wodzicki (2000)

DEBATE: End Member Model I

Slow and Gradual Surface Uplift



Used climate modeling to show that surface uplift of the Andes causes non-linear cooling of land surface and increase in convective rainfall. Argued that gradual surface uplift is permissible with our data.

Surface uplift rates = 0 to 0.25 mm/yr

- MECHANISMS: Crustal shortening, Ablative subduction
- PREDICTIONS: deformation coupled with surface uplift

Ehlers and Poulsen (2009, EPSL); Poulsen et al. (2010, Science)

Sedimentary deposits record δ^{18} O and T of surface water: Used to quantify paleoelevations



Stable isotope paleoelevation estimates:

- Ghosh et al. (2006) ${}^{13}C-{}^{18}O$ clumped isotope thermometer compared paleotemperature and $\delta^{18}O_{mw}$ from authigenic calcite to modern T $\delta^{18}O_{mw}$ vs. altitude gradients
- Garzione et al. (2006) δ^{18} O paleoaltimetry compared $\delta^{18}O_{mw}$ δD_{mw} from authigenic calcite to modern δ^{18} O vs. altitude gradient

Sedimentary deposits record $\delta^{18}O$ and T of surface water: Used to quantify Paleoelevations



δ^{18} O-Altitude Relationship of Meteoric Waters



data from Gonfiantini et al. (2001); Garzione et al. (2007); Bershaw et al. (2010)

Temperature-Altitude Relationship



Temperature corrected at ≤2 km for warmer conditions under lower Andes scenarios (Ehlers and Poulsen, 2009).

Kar et al. (in review); modified from Garzione et al. (2014)

Meteoric Water $\delta^{18}O$

$$\delta^{18}O = \begin{bmatrix} {}^{18}O/{}^{16}O_{\text{sample}} \\ {}^{18}O/{}^{16}O_{\text{reference}} \end{bmatrix} * 1000 \text{ (per mil)}$$

- Sedimentary carbonates: $\delta^{18}O_{carbonate}$ depends on $\delta^{18}O_{meteroric water (mw),}$ evaporation, and temperature of carbonate precipitation
- Temperature dependent fractionation:

 $1000 \ln \alpha_{(Calcite-H20)} = 18.03 (10^{3} T^{-1}) - 32.42$ (Kim and O'Neil, 1997)



Plant Waxes: δD of C₂₅ to C₃₁ n-Alkanes



Rapid Surface Uplift vs. Gradual Surface Uplift



End Member Model II



Rapid Surface Uplift

- Decoupled deformation and surface uplift
- Proposed Mechanisms: Removal of dense lower lithosphere and/or lower crustal flow



Large magnitude surface uplift of the central Eastern Cordillera between 24 and 17 Ma.

Leier et al. (2013, EPSL)

Along-strike variations in surface uplift







~10 Ma Paleosurface:



Gubbels et al. (1993); Kennan et al. (1997); Barke and Lamb (2006); Hoke and Garzione (2008)

Central Andean Plateau



Hoke and Garzione (2008, EPSL)

~45 Ma Paleotopography



Evidence for a long-lived magmatic arc suggests that there was significant relief in the Western Cordillera; stable isotope evidence – Saylor et al. (2014)

~45 – ~25 Ma Paleotopography



- Basin analysis, structural evidence, and thermochronology indicate that the Eastern Cordillera began to deform by ~45 to 40 Ma (Horton et al., 2001; McQuarrie, 2002; Gillis et al., 2006; Barnes et al., 2008)
- Paleoelevation estimates from Salla Fm (Leier et al., 2013) also support this

~20 to ~10 Ma Paleotopography



Leier et al. (2013) results suggest that the Eastern Cordillera experiences a rapid pulse of surface uplift between ~24 and 17 Ma.

<6 Ma Paleotopography

Stable isotope results and incision histories suggest a rapid pulse of surface uplift between 10 and 6 Ma that established the modern topography of the central portion of the Andean plateau.

Along-strike variations in paleotopography

Balance cross sections detailed in Figure 8

Garzione et al. (2017, AREPS)

Along-strike variations in paleotopography

Prediction: Rapid removal of lower lithosphere surface uplift decreases the horizontal compressive stress in the Andean plateau

Observation: Contractional deformation ceased in the central Altiplano and propagated eastward into the Subandes.

Middle-late Miocene events in the Andean plateau

Garzione et al. (2008, Science)

Middle-late Miocene events in the Andean plateau

Garzione et al. (2008, Science)

What mechanisms form broad, high elevation, low relief orogenic plateaus? CAUGHT – Central Andean Uplift and the Geodynamics of High Topography

- Incision history
- Crustal structure
- Shortening history
- Volcanics
- Mantle structure

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Garzione et al. (2017 – AREPS)

The case for convective removal of the lower lithosphere

- Crustal thickening precedes pulses of surface uplift
- Pulses of surface uplift should correspond with an outward jump in deformation
- If high density eclogitic lower crust drives lower lithosphere removal, then crustal thickening history would predict crustal thickness in excess of modern thickness.
- Crustal and mantle structure may show evidence of recent or ongoing convective removal of lower lithosphere.

Along-strike variations in surface uplift

Large magnitude surface uplift of the Altiplano Propagates from S to N from middle to late Miocene time.

Along-strike variations in the outward jump in deformation & aridification of the Altiplano

- Middle Miocene aridification of the southern Altiplano and western slope versus late Miocene aridification of the northern Altiplano
- Middle Miocene onset of southern Subandean deformation versus late Miocene onset of southern Subandean deformation

Lease et al. (2016)

Along-strike variations in crustal shortening

Garzione et al. (2017, AREPS)

Crustal shortening history

Central Central Andean Plateau

Southern Central Andean Plateau

Garzione et al. (2017, AREPS), from McQuarrie et al. (2002, 2008), Barnes et al. (2012)

Crustal thickening versus surface uplift - Central Plateau

Garzione et al. (2017, AREPS), from Eichelberger et al. (2015), Garzione et al. (2006, 2008), Leier et al. (2013)

CAUGHT-PULSE broadband station map

Crustal thickness map (receiver functions) & Crustal/mantle structure

Garzione et al. (2017, AREPS), from Ward et al. (2013, 2016) & Ryan et al. (2016)

The case for lower crustal flow

- Rapid surface uplift during time periods of crustal thickening that is decoupled from crustal shortening
- Modern crustal thickness exceeds that which can be accounted for by crustal shortening
- Trace element evidence for rapid crustal thickening in the absence of shortening
- Exhumation/incision events track both crustal thickening and surface uplift in the absence of shortening

Crustal thickness map (receiver functions) & Crustal/mantle structure

Garzione et al. (2017, AREPS), from Ward et al. (2013, 2016) & Ryan et al. (2016)

Along-strike variations in crustal shortening

Garzione et al. (2017, AREPS)

Crustal shortening history

Central Central Andean Plateau

Garzione et al. (2017, AREPS), from McQuarrie et al. (2002), Barnes et al. (2012), Perez et al. (2016)

Crustal thickening versus surface uplift

Garzione et al. (2017, AREPS), from Eichelberger et al. (2015), Kar et al. (2016), Perez et al. (2016)

front age

-30

Pliocene–recent incision of the NE Andean plateau

Lease and Ehlers (2013)

Summary – Pulsed nature of surface uplift

- The central Central Andean Plateau shows at least two pulses of rapid surface uplift: early Miocene Eastern Cordillera and late Miocene Eastern Cordillera and Altiplano.
- In the Altiplano, rapid surface uplift propagates from south to north from middle Miocene to late Miocene/Pliocene time.

Evidence for lower lithosphere removal & lower crustal flow

- Crustal thickening predictions from balanced cross sections show thickness excess in regions that have experienced multiple pulses of surface uplift → suggests convective removal of eclogitic lower crust
- The northernmost portion of the Andean plateau lacks the magnitude of crustal shortening required to account for the modern crustal thickness → suggests crustal flow into this region
- In the northernmost portion of the Andes, HREE depletion (i.e., high La/Yb) over past ~5 Ma suggests crustal thickening in the absence of crustal shortening over the same time period as surface uplift
 → suggests crustal flow into this region